

Another important result has been obtained from the global reference in the two coast of the strait of Gibraltar. The Real Observatorio de la Armada (ROA) is carrying out this research. One result is that the mean sea level of the station in Ceuta (African coast) with those in Tarifa, Algeciras y Cádiz (European coast) can be compared as it is shown in table 1.

| Station | MSL over TGZ | Altitude of TGZ Related to WGS-84 |
|-----------|--------------|-----------------------------------|
| Cádiz | 2.035 meters | 43.197 meters |
| Tarifa | 1.090 meters | 40.492 meters |
| Algeciras | 0.76 meters | 41.713 meters |
| Ceuta | .887 meters | 41.314 meters |

Table 1. Mean sea level (1990-1999) over TGZ. Altitude: TGZ related to WGS-84 (ROA).

Concerning to the chart datum, the Instituto Hidrografico de la Marina (IHM), is recalculating the hidrographic zero at each stations using the new protocols established by the Bureau Hydrographic (BHO).

2. Strategy

Over the last years, with the use of the new technologies, new type of system measurement equipment has been developed. The Spanish network is using a modern system: Acoustic gauges. Also, for the traditional float, an encoder and a data logger with data modem transmission facility has been incorporated, but also a tide gauge with radar technology is being testing.

But the most important advance for the sea level monitoring is the modern geodetic techniques for geodetic fixing of the tide gauge bench mark. With this technique it will be possible not only to calculate the absolute sea level by removing the crustal movement but also to refer the sea level measurement to a global reference system. Then the tide gauge sea level data can be comparable with the altimetry data and therefore a better knowledge of the spatial variation will be obtain. In Spain, as mention before, this technology is being used in many stations. The strategy for the future is to select some station for sea level long-term evolution. In those stations a permanent GPS will be install. At each of other stations, an annual or biannual GPS campaign will be carry out.

Concerning to the data analysis, a preliminary trend of the relative monthly mean sea level time series are presented in figure 3.

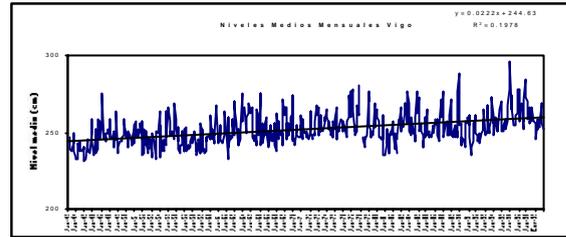


Figure 3. Vigo Mean Sea Level Evolution

But in the future, applying a mathematical model to the monthly mean sea levels in order to remove the meteorological effects will be perform the determination of the long-term mean sea level and its evolution. The method, as described in (Pugh, 1987), consists of fitting a mathematical function to the time series of $Z_0(t)$ by a least square technique. The analytical expression to include in the model is:

$$Z_0(t) = \bar{Z}_0(t) + at \quad (\text{long-term mean and trend})$$

$$+ N + S_a + S_{aa} \quad (\text{periodic tidal terms})$$

$$+ b_0 P_A + b_1 P_A \setminus x + b_2 P_A \setminus y \quad (\text{meteorological effects})$$

3. Data processing

The sea level is measured in 1, 5 or 10 minute's intervals. It varies in function of the different purposes. For example to study the seiches that appears frecuently in Palma de Mallorca, the data are registering every minute. The raw data are transmitted to the Data Center and a quality control procedure is performed to eliminate obvious errors as clears spikes, date and changes in references.

Then, clean data are filtered to obtain the hourly data. The two filters mainly used are the 54 points symmetrical filter (Pugh, 1997) and the consecutive mobile means $A_m A_m A_n$ (Godin, 1972). The results of these two filters are different and when the Godin filter is used the lost energy should be recover at harmonic analysis step.

The hourly data are processing with the software developed by the University of Hawaii Sea level Center (Caldwell, 1998) to perform; harmonic analysis and prediction (Foreman, 1997), hourly quality control by inspection of the residual, and, the daily mean values by removing the diurnal and semidiurnal tidal constant.

The monthly and annual mean sea level is calculated by a simple weight mean. The tidal ranges are calculated on an annual basis.

4. Annual Mean Sea Level Evolution.

In this paper, the values obtained for the mean sea level are the relative mean sea level in which there are included not only the crustal vertical movement but also the meteorological effects. The information about the crustal movements is very poor. In fact the results presented by (Emery & Aubry, 1991) are based in the relative mean sea level trends.

The annual mean sea level evolution during the decade (1990-1999) shows in figure 4 an increase sea level from 1992 to 1997. This year presents the higher level in almost all of the station. But also it can be seen that in the Sevilla-exclusa station, located in the river Guadalquivir, the extremely high level of this year give an idea of the large impact of the storms. As a complementary information, in the pacific the sea level rise 25 cm as the impact of the Niño 1997. In Spain during the decade 1980-1999 (Lavin&García ,1992) the highest mean sea level occurs in 1983 and 1987, the same years as the Niño 1982-1983 and 1987.

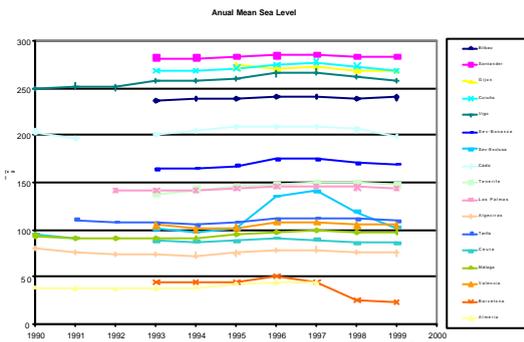


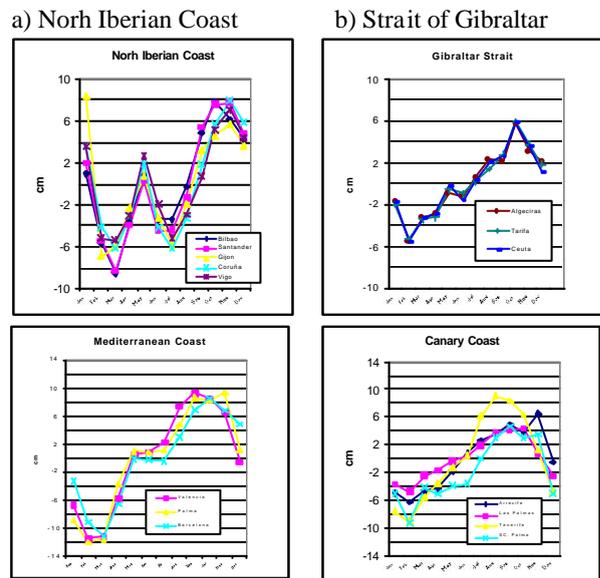
Figure 4. Anual mean sea level.

5. Seasonal variability.

The seasonal variability has a close relation with the wind and atmospheric pressure. Then, as the monthly data are not removed from this effect, a marked seasonally in monthly average appears. The behavior of the seasonal variability is presented in figure 5a-5d the results are grouped in regions according to the climatology of each region. In the Iberian Peninsular Coast the seasonal variability shows a clear decrease of sea level in the summer time, and an increase in the autumn season. Lowest values occurs in July or august and higher in October or November. The range is about 16 cm. (Lavin & García, 1992) and has been found a clear relation of this variability with the atmospheric pressure, high barometric factors, and also with the wind stress.

The seasonal variability in the Strait of Gibraltar and in the Mediterranean shows a gradual increase in the sea level between February and October. Then it begins to decreases until February. The ranges are about 12 in the Gibraltar Strait and about 22 in the Levantine Coast. The main causes of this observer variability are the atmospheric pressure, the wind and the volume thermal expansion. In summer the atmospheric pressure is fairly low in this region (less than 1020 mb). On the other hand, in this area the wind blows during the summer from East to West. So the East wind flowing in the summer months piles up the water in the Spanish Mediterranean coast. Furdermore the sea temperature is rather warmer in summer than in winter (Medatlas climatology, 1996).

The seasonal variability in Canary Island presents its minimum in February and the maximum appears in August or September instead of October-November. The rage of the variability is around 10 cm in all the stations except in Tenerife (20 cm). Part of the progressive sea level increase between June and August could be explained by the onset of the trade winds, which piles up the water in the eastern of the island. ISEMER and Hasse (1987) shows highest trade winds stress forces during the summer, about $.12N/m^2$, twice the January value. The seasonal variability in S.C de la Palma and in Arrecife have been calculated taking into account the last 2-3 last years. For this reason, the pick that appears in November could be due to a particular behavior of the sea level during those years.



c) Mediterranean Coast d) Canary Island

Figure 5. Seasonal Variability.

6. Tides

The Harmonic analysis (Foreman 1977) has been applied to the hourly data for periods of 366 days to

obtain the main constituents. The semidiurnal constituents are dominant (Form Factor < .25) in all the atlantic coast. In the Mediterranean Coast, Malaga has also a semidiurnal tide but it is close to a mixed tide (form factor = 0.21) but Valencia presents a mixed tide, mainly diurnal (Factor Form=2.67) and became mainly semidiurnal in Barcelona (Factor Form = 0.67).

The mean values of the M2 and its standard deviations have been obtained using the years, within the decade, that have enough data and quality, to perform the harmonic analysis. The standard deviations of the amplitudes are less than 0.2 cm, larger in Coruña and Cádiz. The standard deviations of the phase are about 0.5 (g) and much large in some stations about 1.1(g), but probably, is because of some time error on the data. In the figure 8 the amplitude and phase of the M2 constituent shows its propagation along the coast from south to north in the Atlantic coast coast.

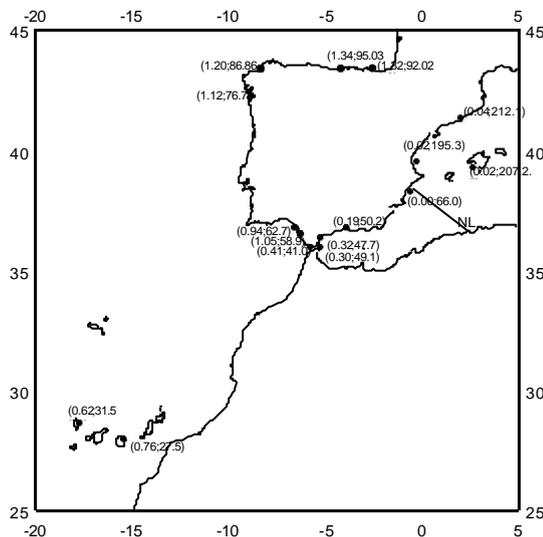


Figure 8: M2. Amplitude and Phase

The Strait of Gibraltar is a very complex area and more gauges would be necessary to describe with detail the propagation of the tide in this region. A very complete description can be found in Garcia Lafuente (1996), who used a network composed of 4 temporary pressure gauges at each side of the strait. One of the most important features is that the M2 amplitude increase from East to West along the strait. In the figure 2, it can be seen that the amplitude of tarifa is 9 cms higher than in Algeciras.

In the Mediterranean Coast, the amplitude of M2 is increasing from Alicante to Algeciras and in the opposite direction from Alicante to Barcelona, and the big jump of the phase between Alicante and Valencia shows the existence of the well known Alicante–Oran nodal line.

The maximum and minimum tidal ranges found for the decade is presented in table 2. that shows that the tides in the Mediterranean are very small and in the Atlantic the maximum ranges decreases from north to South and there are a big differences in the spring tides and neap tides

| Station | Tide range (cm) | | Station | Tide range (cm) | |
|-----------|-----------------|-----|------------------|-----------------|-----|
| | Max | Min | | Max | Min |
| Bilbao | 503 | 117 | Ceuta | 119 | 18 |
| Santander | 501 | 118 | Algeciras | 122 | 16 |
| Coruña | 458 | 103 | Málaga | 84 | 9 |
| Vigo | 421 | 99 | P. Mallorca | 23 | 0 |
| Bonanza | 347 | 82 | Las Palmas | 297 | 50 |
| Cádiz | 407 | 76 | S.C. de la Palma | 244 | 44 |
| Tarifa | 163 | 26 | | | |

Table 2. Maximum and minimum tide ranges.

Conclusion:

The sea level is monitoring in Spain taking into account the new technology not only in the measurement equipment system but also in the geodetic fixing. This way of research will improve to understand the long-term evolution of the mean sea level and the spatial variation.

The annual mean sea level evolution during the decade shows the higher level in 1997 in all the stations and also appears an increasing trend from 1992 to 1997. However, this result seems to be influenced by the climatological cycles. Anyway the determination of the mean sea level and the long-term trend need to be performed by removing the atmospheric influence.

The seasonal variability is higher in the Mediterranean (22cm) than in the Atlantic (10cm-12cm) with the exception of Tenerife.

The tide ranges are very low in the Spanish Mediterranean Coast.

References

Caldwell, P. C. 1998. Sea Level Data Processing On IBM-PC Compatible Computers. Version 3.0 (Year 2000 Compliant). JIMAR Contribution N°98-319.

Emery, K.O, Aubrey, D.G (1991) Sea Levels, Land Levels, and Tide Gauges. Springer-Verlag, 237 pp.

Foreman, M.G. 1977. Manual of tidal height analysis and predictions. Pacific Marine Science report. (Institute of Ocean Science Patricia Bay) 77(10): 101 pp.

García Lafuente, J. (1986) Variabilidad del nivel del mar en el Estrecho de Gibraltar. Mareas y oscilaciones residuales. Phd. Thesis, 153 pp.

Godin, G. 1992. The analysis of tides. University of Toronto Press. Toronto: 264 pp.

IOC. Manual on sea-Level Measurement and Interpretation. Volume II- Emerging Technologies. Manual and Guides 14.

ISEMER, H & Hasse I. (1987) The bunker climate atlas of the North Atlantic Ocean, 2, Air Sea Interaction, 252 pp.

Lavin, A. & García, M.J. (1992). Mean Sea level along the North Atlantic Spanish coast 1980-1989. ICES mar. Sci.Symp, 195, 187-192.

Mediterranean Hydrological Atlas, 1996. A Mast Supporting Initiative. MAS2-Ct93-0774.

Perez B., García M.J., Puyol B. (1999) Integración y Optimización de las Redes Mareográficas Españolas (in press). Presented to the Jornadas Portuarias, Coruña 1999.

Pugh, D.T. (1987) Tides, surges and mean sea level. Chichester: John Wiley & sons, 472 pp.